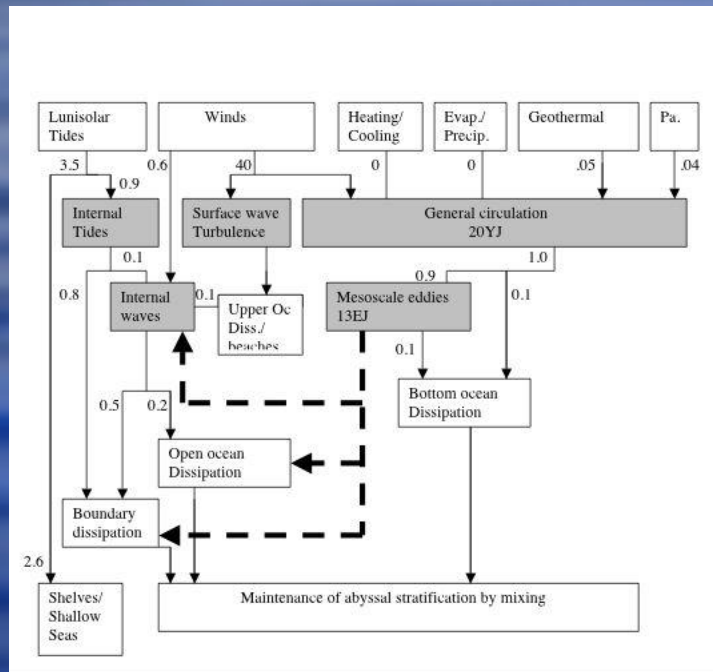


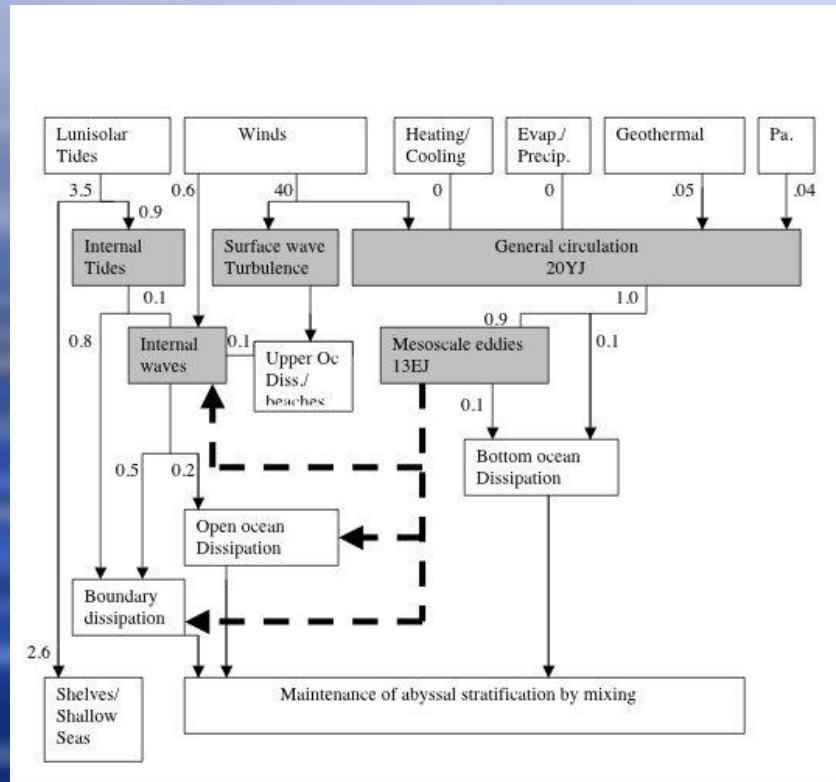
Inviscid Dissipation in Ocean Climate Models

FCI Seminar

March 24, 2011

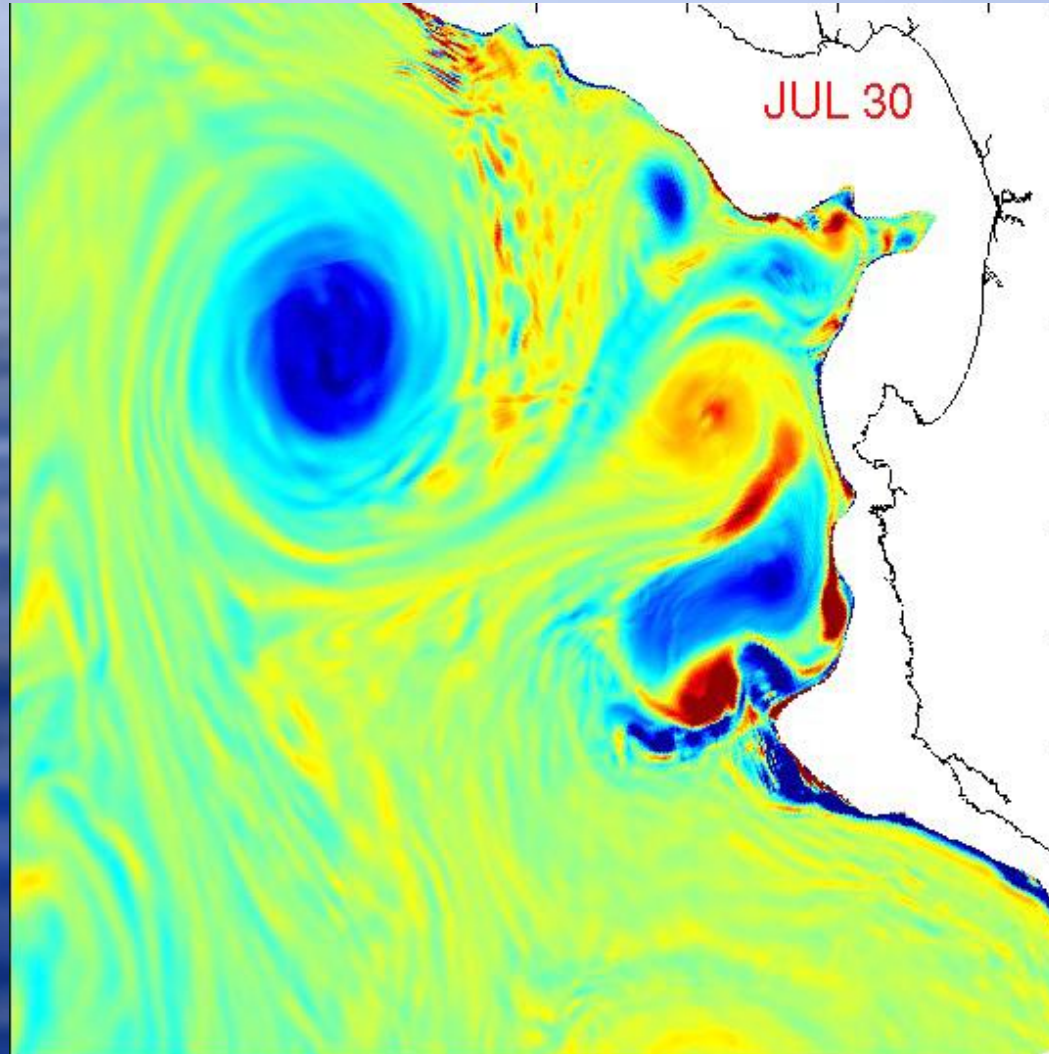


An analysis of ocean power, after Wunsch and Ferrari, 2004.

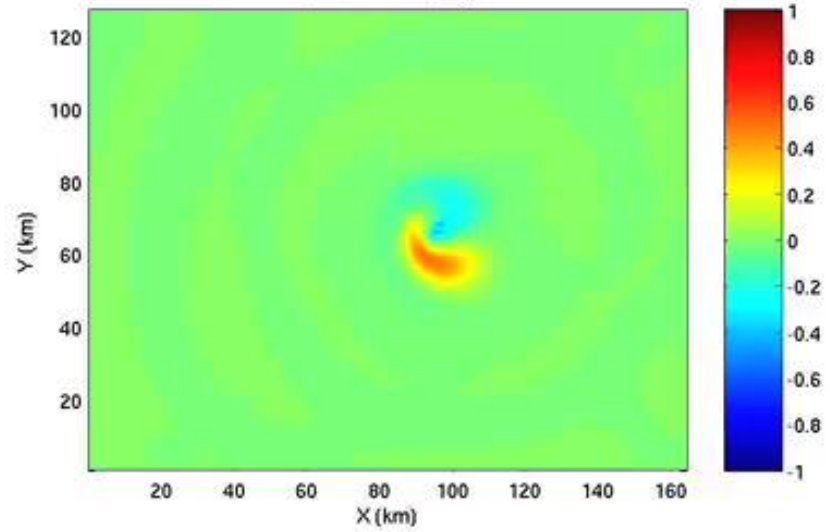


And there is a related potential vorticity budget question.

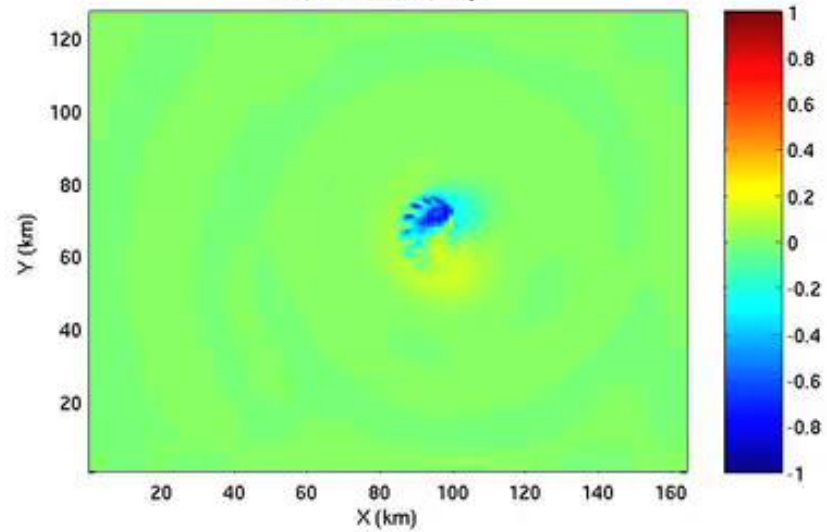
Numerical Evidence in Support An ultra-fine embedded solution for Monterey Bay



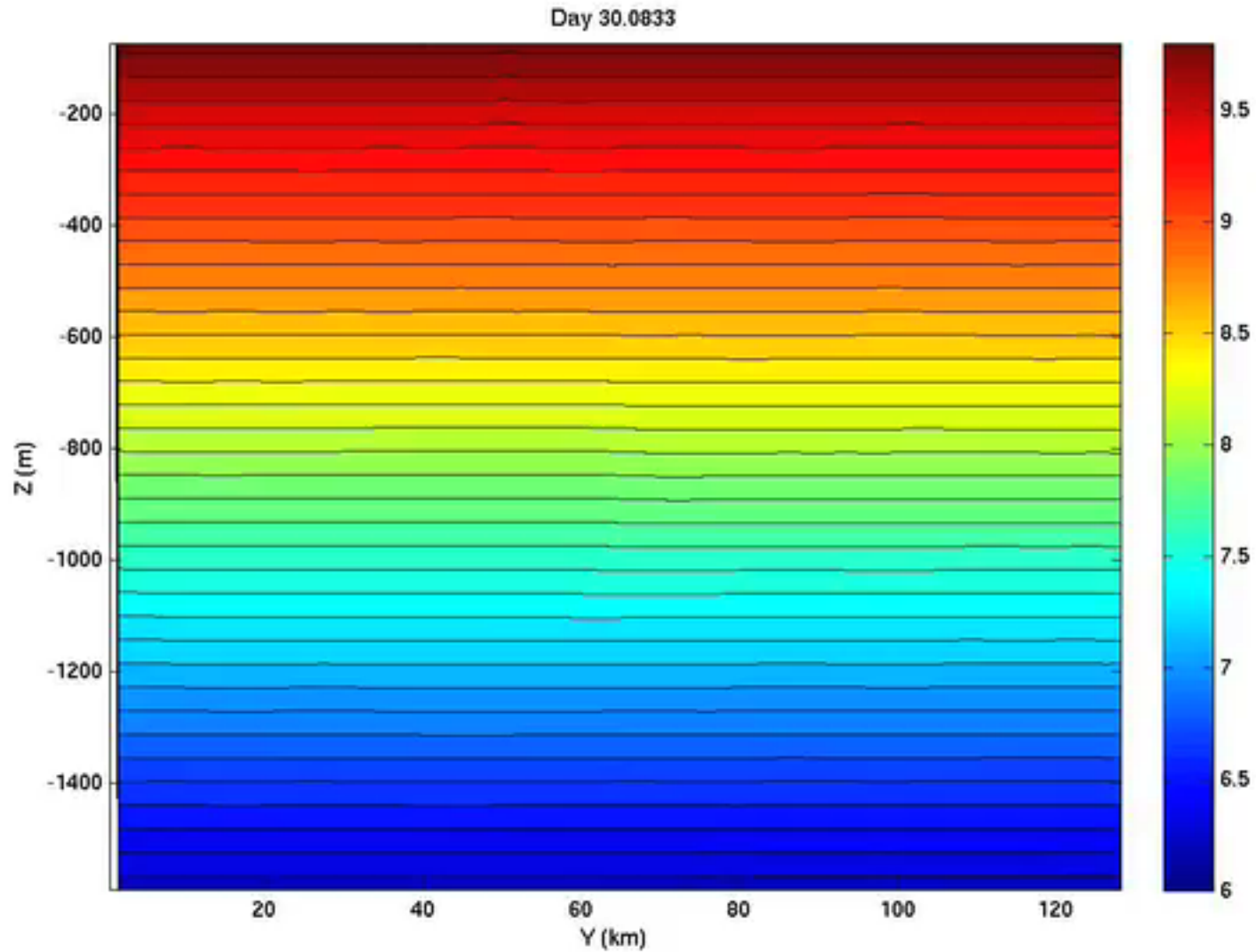
Ro; Z = -713.169; Day 1



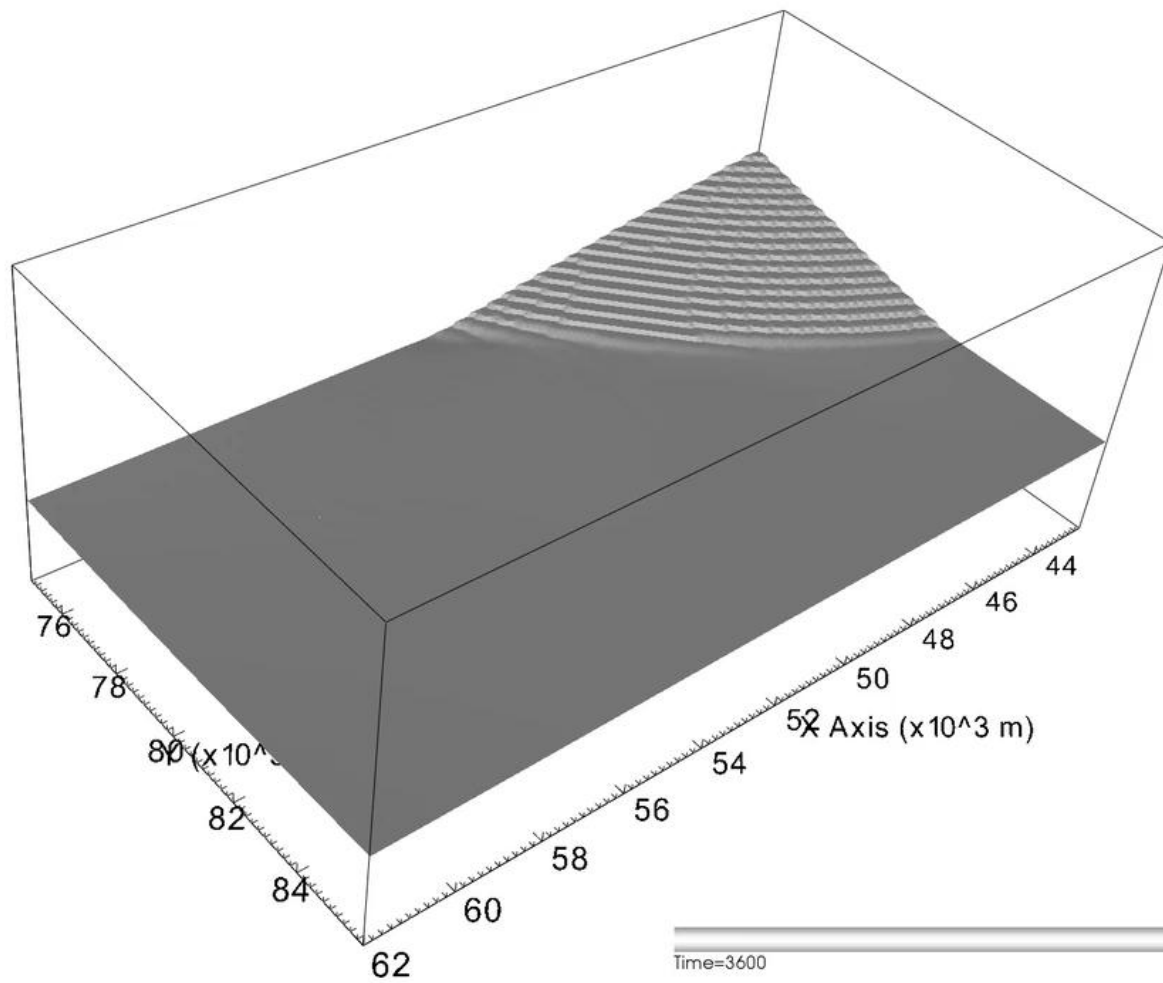
Ro; Z = -988.689; Day 1



Temperature at western wall



Contour
Var: Temp
7.400
Max: 9.840
Min: 0.000



Summary:

Something very interesting happens at topography when balanced flow interacts with it.

Question:



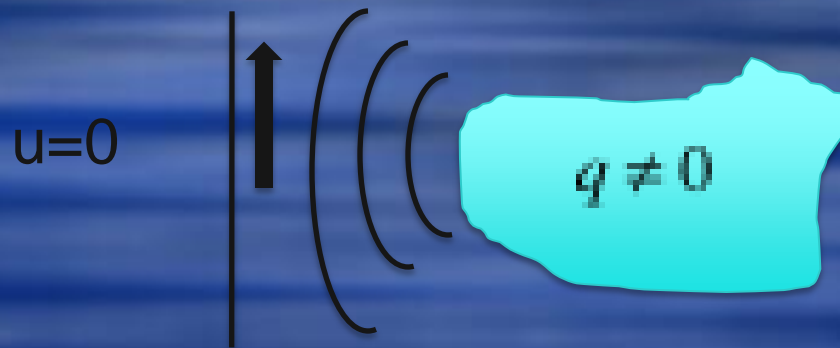
At the wall, normal flow vanishes
and consider inviscid, adiabatic flow

$$\cancel{u}_t + \cancel{u}u_x + \cancel{v}v_y - f\cancel{v} = -M_x$$

$$v_t + \cancel{u}v_x + \cancel{v}v_y + \cancel{f}u = -M_y$$

$$q_t + \cancel{u}q_x + \cancel{v}q_y = 0; \quad q = (f + v_x - \cancel{u}_y) / z_\rho$$

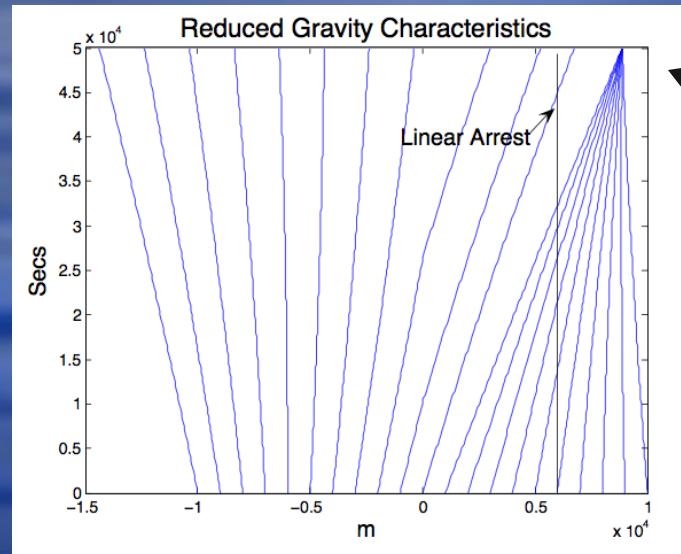
$$M_{xx} - f^2 \frac{M_{\rho\rho}}{M_{\rho\rho}} = 0$$



What does nonlinearity do?

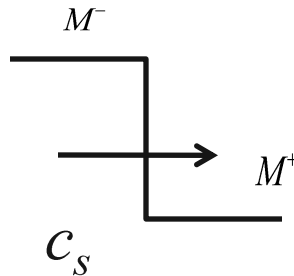
$$v_t + \left(\frac{v^2}{2} \right)_y + (v_x v)_y + M_y = -v_{yy} - v_x v_{xy} - M_{xy}$$

Shock Formation (speed can be computed)



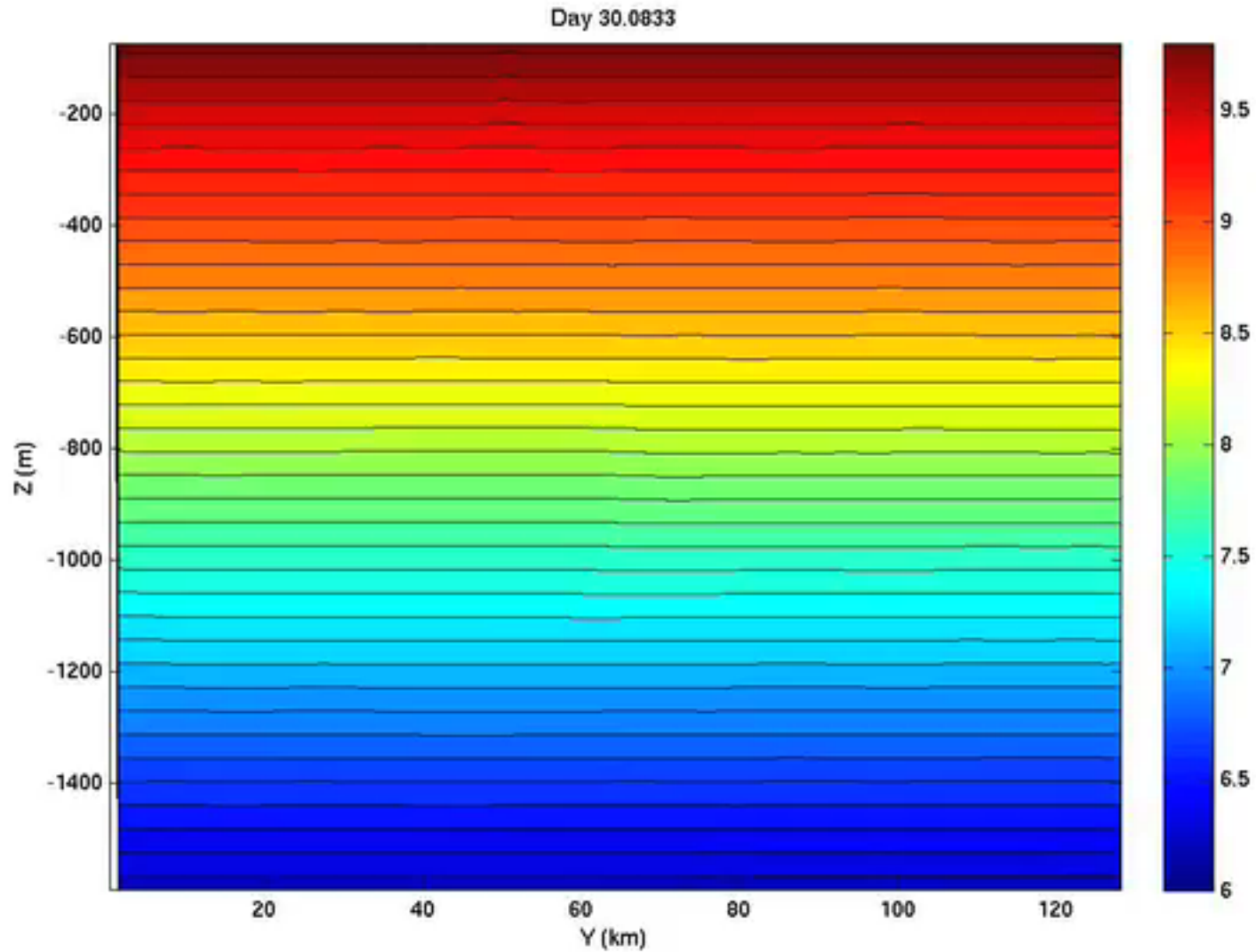
How fast does the shock go?

$$-c_s v_y + \left(\frac{v^2}{2} \right)_y + (v_g v)_y + M_y = (F(v_g))_y$$



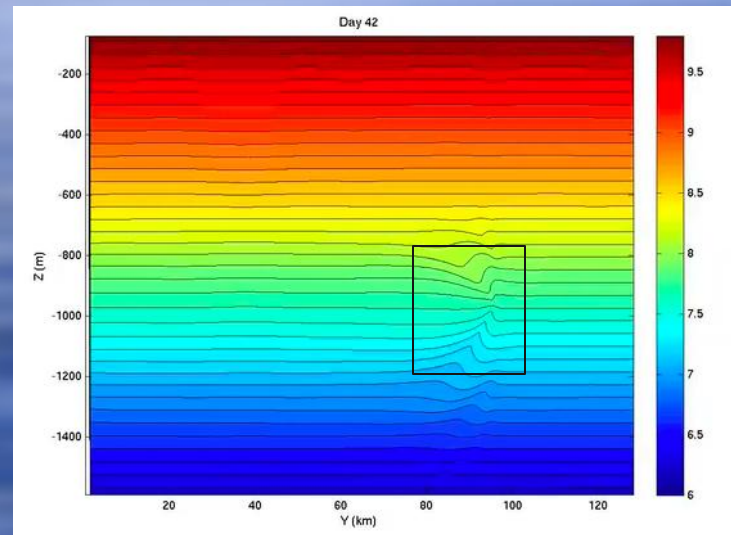
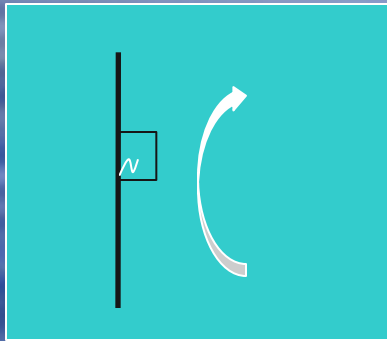
$$c_s = \frac{M^+ + M^-}{2 fR} + v_g - c - fR$$

Temperature at western wall



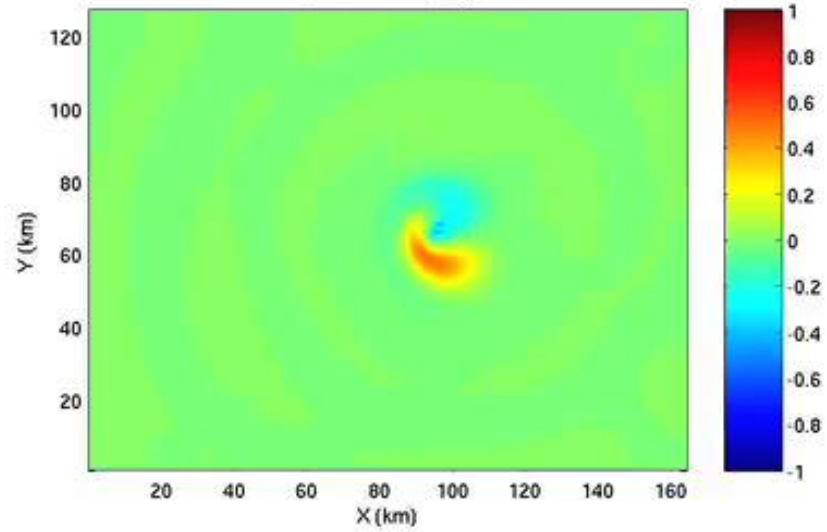
Non-hydrostatic dynamics arrest the shock

$$(z_\rho q)_t + (uqz_\rho)_x + (vqz_\rho)_y = (Y_x - X_y) + (u_\rho H)_y - (v_\rho H)_x$$

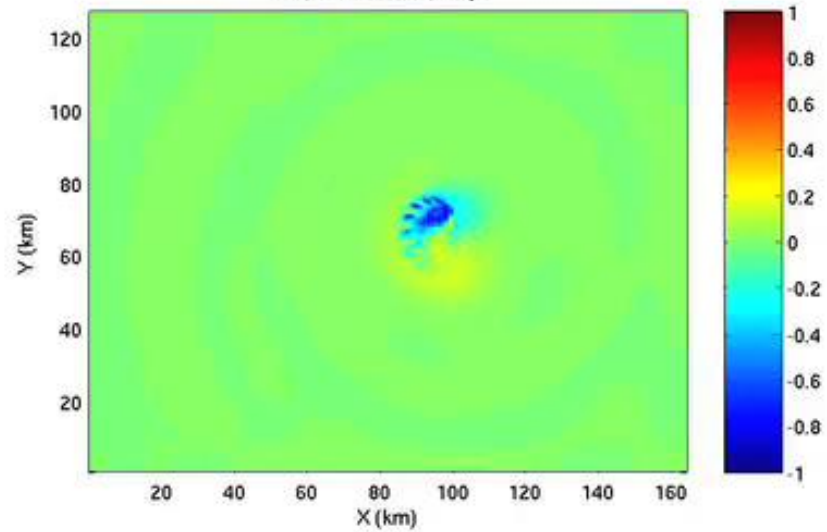


$$\iint uz_\rho q' dy d\rho = \iint [v_\rho H]_o dy d\rho = O(fR)^2$$

Ro; Z = -713.169; Day 1



Ro; Z = -988.689; Day 1



Summary and to do list

Simple, surprisingly accurate local theory for wall interactions can be written

- steepening and breaking
- can be parameterized

Extend this to non-wall topography?
(Yes)?

Introduce into models to control boundary dissipation